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Al-free di-trioctahedral substitution in chlorite and a ferri-sudoite end-member

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ABSTRACT: A compilation of Fe^{3^+} -bearing chlorite analyses is used: (1) to investigate the Alfree di-trioctahedral (AFDT) substitution $2\text{Fe}^{3^+} + \Box = 3(\text{Mg}, \text{Fe}^{2^+})$ in chlorite; and (2) to estimate the composition of a ferri-sudoite end-member $(\text{Si}_3\text{Al})[(\text{Fe}^{2^+}, \text{Mg})_2\text{Fe}_2^{3^+} \Box \text{Al}]O_{10}(\text{OH})_8$ within the chlorite solid-solution domain. According to our observations, up to two Fe^{3^+} cations might be allocated in the M2-M3 chlorite sites by the substitution AFDT, which does not involve Al. These unexpected observations were made possible by the development of μ XANES techniques allowing *in situ* measurements of $X\text{Fe}^{3^+}$ ($\text{Fe}^{3^+}/(\text{Fe}^{2^+} + \text{Fe}^{3^+})$) in heterogeneous chlorite. Although further studies are required to confirm the crystallographic position of Fe^{3^+} and refine its ionic/ magnetic behaviour in chlorite, this development creates opportunities for developing new geothermometers.

KEYWORDS: chlorite, µXANES, geothermometry, sudoite.

Chlorite is a ubiquitous phyllosilicate composed of tetrahedral and octahedral layers according to the 14.2 Å 2:1:1 arrangement (*e.g.* Meunier, 2005). The tetrahedral sheet hosts tri- or tetravalent cations (Si⁴⁺, Al³⁺ or rarely Ti⁴⁺, Fe³⁺), while the octahedral one is composed of di- or trivalent cations (usually Fe²⁺, Mg²⁺, Mn²⁺, Al³⁺ and Fe³⁺). The general formula of chlorite is: $R_{6-x-3y}^{2+}R_{x+2y}^{3+} \square_y$ (Si_{4-x} R_x^{3+})O₁₀(OH)₈, in which R^{2+} represents divalent cations, R^{3+} trivalent cations and \square octahedral vacancies (*e.g.* Wiewióra & Weiss, 1990). According to Bailey (1988), the chlorite structure is formed of two types of tetrahedral crystallographic sites (denoted T1 and T2) and two octahedral sites (M3 and M4) are located in the brucite

interlayer space. AlVI or trivalent Fe3+ cations are located preferentially in the M4 site, while the divalent cations occupy the other sites (Bailey, 1988). There are three categories of chlorite in diagenetic and low- to high-grade metamorphic rocks: tri-trioctahedral (sum of octahedral cations close to 6 a.p.f.u), di-trioctahedral (5 a.p.f.u) and di-dioctahedral (4 a.p.f.u). The main chemical substitutions occurring in chlorite are: Tschermak (TK) $Si_{(IV)}^{4+}$ + $(Fe^{2+}, Mg^{2+})_{(VI)} = Al_{(IV)}^{3+} + Al_{(VI)}^{3+}$; ferromagnesian (FM) $Mg_{(VD)}^{2+} = Fe_{(VD)}^{2+}$ and di-trioctahedral (DT) $2Al_{(VD)}^{3+} + \Box_{(VD)}$ $=3(Mg^{2+},Fe^{2+})_{(VD)}$. Two filling models are used in the literature for the distribution of cations on the crystallographic sites. The first model referred to as 'disorderly', considers that there is no preferred configuration (e.g. Inoue et al., 2009). The second model, referred to as 'ordered' is constrained by specific rules for filling cations in the crystallographic sites such as preferred configuration and Fe-Mg equipartition (e.g. Vidal et al., 2001, 2005, 2006; Lanari et al., 2014a). These authors assume, following Holland et al. (1998), Al_{IV} in T2, vacancies in M1, Al_{VI}

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	T1(2)	T2(2)	M1(1)	M2-M3(4)	M4(1)
Amesite-Mg/corundophilite	Si,Si	Al,Al	Al	Mg,Mg,Mg,Mg	Al
Amesite-Fe	Si,Si	Al,Al	Al	Fe,Fe,Fe,Fe	Al
Clinochlore	Si,Si	Si,Al	Mg	Mg,Mg,Mg,Mg	Al
Chamosite/daphnite	Si,Si	Si,Al	Fe	Fe,Fe,Fe,Fe	Al
Sudoite-Mg	Si,Si	Si,Al		Mg,Mg,Al,Al	Al
Sudoite-Fe	Si,Si	Si,Al		Fe,Fe,Al,Al	Al
Al-free chlorite	Si,Si	Si,Si	Mg	Mg,Mg,Mg,Mg	Mg
Pyrophyllite-gibbsite	Si,Si	Si,Si		Al,Al,Al,	Al
Ferri-sudoite	Si,Si	Si,Al		$(Fe,Mg)_2,Fe^{3+},Fe^{3+}$	Al

TABLE 1. Structural formulae for classic chlorite end-members and the new ferri-sudoite end-member.

The numbers in parentheses refer to the quantity of atoms (per structural formula) distributed in tetrahedral (T) or octahedral (M) sites.

in M4 then M1 and then, eventually, in M2-M3. Fe and Mg in M2-M3 then in M1. Following the equipartition assumption, XMg $(Mg^{2+}/(Mg^{2+} + Fe^{2+}))$ is considered to be equal between the M1 and M2-M3 sites. Several chlorite end-members are classically defined (Table 1): Al-free chlorite (Si₄)[Mg₆]O₁₀(OH)₈, amesite (Si₂Al₂) $[(Fe^{2+},Mg)_4Al_2]O_{10}(OH)_8,$ clinochlore (Si₂Al) [Mg₅Al)O₁₀ (OH)₈, corundophilite (Si₂Al₂)[Mg₄Al₂] $O_{10}(OH)_8$, daphnite $(Si_3Al)[Fe_5^{2+}Al]O_{10}(OH)_8$, pyrophyllite-gibbsite (Si₄)[Al₄]₂]O₁₀(OH)₈ and sudoite $(Si_3Al)[(Fe^{2+},Mg)_2\squareAl_3]O_{10}(OH)_8.$ These endmembers do not include any ferric iron component.

For 15 years, many authors have emphasized the importance of Fe^{3+} in chlorite, particularly in terms of its use as a geothermometer (Laird, 1988; Vidal *et al.*, 2006; Inoue *et al.*, 2009; Bourdelle *et al.*, 2013; Lanari *et al.*, 2014a). An additional substitution: Al-Fe³⁺ that is supposed to occur on the M4 site was defined but without defining a ferric chlorite end-member (Vidal *et al.*, 2006). In this model, the incorporation of Fe³⁺ changes the chlorite composition and the activity of the other end-members. This choice of the M4 site is based on the results of Smyth *et al.* (1997) suggesting that Fe³⁺ in chlorite shows a strong preference for the M4 site where it substitutes for Al³⁺.

Several methods are available to measure Fe^{3+} in chlorite: chemical titration (Tschermak, 1891; Orcel, 1927; Hallimond *et al.*, 1939; Brindley, 1951), Mössbauer (Pal *et al.*, 1993; Aja & Dyar, 2000; Inoue *et al.*, 2009), X-ray photoelectron spectroscopy (XPS) (Inoue *et al.*, 2010) and micro-X-ray absorption near-edge spectroscopy (μ -XANES) (Wilke *et al.*, *al.*, 2009), Wilke *et al.*, 2009), Wilke *et al.*, 2010, and micro-X-ray absorption near-edge spectroscopy (μ -XANES) (Wilke *et al.*, *al.*, 2010).

2001; Vidal et al., 2006; Rigault, 2010; Trincal et al., 2015).

CHLORITE DATA FROM THE LITERATURE

A compilation of almost 200 published data of chlorite structural formulae containing Fe³⁺ is reported in Table 2. These chlorites come from a wide range of rocks from various localities. Some are of detrital origin; another experienced diagenetic to metamorphic conditions as indicated by the oscillatory zoned chlorites from the Pic-de-Port-Vieux outcrop (Fig. 1; Trincal et al., 2015). In addition, 202 data from three localities were incorporated in this review (Inoue et al., 2009). Most of the chlorites selected contain (in a.p.f.u. see Table 2): 2.5 to 3.5 Si⁴⁺, 2 to 4 Al³⁺ (Fig. 2), 1.5 to 5.5 R^{2+} (Mg²⁺, Mn²⁺, Fe²⁺) (Fig. 3) and <1 Fe³⁺ (Fig. 4). These analyses are distributed between classical end-members and apparently follow the classical DT and TK substitutions (Figs 2, 3) or the Al-Fe $^{3+}$ substitution (Fig. 4).

In diagenetic to metamorphic rocks, the XFe^{3+} content of chlorite depends on the temperature and the oxygen fugacity that is controlled by the buffering assemblage, *i.e.* the bulk-rock composition (Lanari *et al.* 2014a). This control is supported by data from Inoue *et al.* (2009), which show relationships between sampling localities and chlorite structural formulae and oxidation state (Figs 2–4). It is therefore critical to look at the variability of XFe^{3+} recorded by chlorite in one rock sample, for a given bulk-rock composition. In such cases the observed variability of XFe^{3+} is caused by temperature variations,

N°	Reference	Method	SiO_2	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	FeO	MgO	MnO	$CaO + Na_2O + K_2O$	Others	Total	Si Ti	^{IV} Al	^{VI} Al	Fe ³⁺	Fe ²⁺	Mg	Mn	
1	Trincal et al. (2015)	XAN	25.17	0.01	22.76		20.98	17.27	0.09	0.05	0.05	86.37	2.58	1.42	1.33	0.68	1.04	2.64		0.30
2	Trincal et al. (2015)	XAN	26.11	0.01	22.93		18.43	19.64	0.09	0.04	0.05	87.29	2.63	1.37	1.35	0.38	1.13	2.95		0.18
3	Trincal et al. (2015)	XAN	25.74	0.01	23.25		20.43	17.86	0.09	0.04	0.05	87.47	2.60	1.40	1.37	0.53	1.14	2.69		0.25
4	Trincal et al. (2015)	XAN	25.98	0.01	23.65		18.39	19.09	0.09	0.04	0.05	87.30	2.62	1.38	1.42	0.32	1.19	2.87		0.18
5	Trincal et al. (2015)	XAN	25.53	0.01	23.16		21.94	16.41	0.08	0.04	0.04	87.23	2.59	1.41	1.37	0.74	1.04	2.48		0.35
6	Trincal et al. (2015)	XAN	25.76	0.02	22.91		21.92	16.36	0.10	0.05	0.05	87.17	2.62	1.38	1.37	0.70	1.08	2.48		0.35
7	Trincal et al. (2015)	XAN	26.39	0.01	23.00		17.59	19.83	0.11	0.05	0.05	87.02	2.65	1.35	1.38	0.36	1.08	2.97		0.20
8	Trincal et al. (2015)	XAN	25.75	0.02	23.17		20.47	17.44	0.11	0.05	0.05	87.04	2.62	1.38	1.39	0.53	1.15	2.64		0.27
9	Trincal et al. (2015)	XAN	25.85	0.01	23.55		18.70	18.83	0.11	0.05	0.05	87.15	2.61	1.39	1.42	0.33	1.22	2.84		0.18
10	Trincal et al. (2015)	XAN	25.55	0.02	23.05		21.48	16.58	0.11	0.05	0.05	86.87	2.60	1.40	1.37	0.72	1.02	2.52		0.35
11	Debret et al. (2015)	XAN	33.76		12.69		4.17	33.65			1.59	86.14	3.23	0.77	0.66	0.20	0.14	4.80		0.04
12	Debret et al. (2015)	XAN	33.73		13.54		3.35	33.92			1.40	87.00	3.19	0.81	0.70	0.17	0.16	4.79		0.03
13	Debret et al. (2015)	XAN	33.47		13.24		3.26	33.66			1.76	85.57	3.21	0.79	0.71	0.13	0.13	4.81		0.02
14	Zhang et al. (2011)	CHE											2.67 0.01	1.33	1.24	0.47	1.51	2.75	0.01	0.19
15	Inoue et al. (2010)	XPS	29.98		15.56	6.36	16.45	16.68	0.40	0.33		85.75	3.13	0.87	1.05	0.50	1.44	2.60	0.04	0.33
16	Inoue et al. (2010)	XPS	28.08	0.07	18.34	3.80	13.82	19.08	1.12	0.31		84.60	2.93 0.01	1.07	1.18	0.30	1.21	2.97	0.10	0.19
17	Inoue et al. (2010)	XPS	27.34	0.03	18.58	4.51	16.42	17.75	0.35	0.27		85.24	2.87 0.00	1.13	1.17	0.36	1.44	2.78	0.03	0.18
18	Inoue et al. (2010)	XPS	29.98	0.04	19.05	3.91	11.26	20.30	0.60	0.29		84.45	3.03 0.00	0.97	1.30	0.30	0.95	3.06	0.05	0.30
19	Inoue et al. (2010)	XPS	24.26	0.01	17.64	5.56	33.82	1.37	3.34	0.11		86.09	2.85 0.00	1.15	1.29	0.49	3.32	0.24	0.33	0.31
20	Inoue et al. (2010)	XPS	27.17	0.05	18.77	3.76	19.37	13.71	3.21	0.24		86.29	2.89 0.00	1.11	1.25	0.30	1.73	2.18	0.29	0.21
21	Inoue et al. (2010)	XPS	22.86	0.03	17.55	6.40	30.55	1.92	2.15	0.13		81.59	2.80 0.00	1.20	1.33	0.59	3.13	0.35	0.22	0.35
22	Rigault (2010)	XAN											2.63 0.01	1.36	1.41	0.13	2.51	1.84	0.01	0.08
23	Rigault (2010)	XAN											2.71	1.29	1.60	0.14	2.58	1.40	0.05	0.22
24	Rigault (2010)	XAN											2.92	1.08	1.62	0.10	1.82	2.11	0.02	0.31
25	Rigault (2010)	XAN											2.86	1.14	1.44	0.34	0.34	3.54		0.32
26	Rigault (2010)	XAN											2.78 0.01	1.21	1.10	0.79	0.79	2.94	0.03	0.34
27	Rigault (2010)	XAN											2.57 0.01	1.42	1.29	0.75	1.84	1.79	0.01	0.30
28	Rigault (2010)	XAN											2.93	1.07	1.53	0.18	0.39	3.57		0.32
29	Singer <i>et al.</i> (2009)	EXA											2.67 0.03	1.31	0.33	1.74	1.51	2.22	0.01	0.38
30	Vidal et al. (2006)	XAN											2.76	1.24	1.37	0.49	2.14	1.62		0.31
31	Vidal et al. (2006)	XAN											2.72	1.28	1.32	0.43	2.11	1.85		0.24
32	Vidal et al. (2006)	XAN											2.71	1.29	1.31	0.14	2.58	1.84		0.07
																		(0	contin	ued)

TABLE 2. Compilation of structural formulae for Fe^{3+} -bearing chlorites, calculated on the basis of O = 14.

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TABLE 2. (contd.)

N°	Reference	Method	SiO ₂	TiO ₂	Al ₂ O ₂	Fe ₂ O ₂	FeO	MgO	MnO	$CaO + Na_2O + K_2O$	Others	Total	Si	Ti	^{IV} A1	^{VI} A1	Fe ³⁺	Fe ²⁺	Mg	Mn	
22	Vidal at al. (2006)	VAN	- 2	- 2	2-3	2-3		0 -		2 -			2 60		1 2 1	1 2 1	0.14	2.60	1 02		0.07
22	Vidal $el al. (2006)$	MÖS	20.70	0.25	22.00	1 42		21.10	0.01	<0.20	0.01		2.09		1.51	1.51	0.14	2.00	1.02		0.07
25	Zazzi $el ul. (2000)$	MÖS	22.60	0.23	12 70	5.05		22.00	0.01	<0.20	0.01		2.00		0.76	0.79	0.01	0.10	4.55		0.15
36	Brandt <i>et al.</i> (2000)	CHE	26.15	0.11	21 13	5.95	17.20	17 73	0.15	<0.74	0.01	88 14	2.67		1 33	1 21	0.15	1 47	4.75		0.07
37	Brandt et al. (2003)	CHE	26.83		21.15	6.07	13.30	20.67				88 51	2.07		1.33	1.21	0.40	1.11	3.07		0.17
38	A ia & Dvar (2002)	MÖS	20.05	0.04	23.03	0.07	37.94	20.07 8 11	0.20	0.05	0.01	00.51	2.07		1.55	1.21	0.45	2 72	1 15	0.03	0.17
39	Aja & Dyar (2002)	MÖS	33.90	0.04	23.75		833	33.80	0.27	0.10	0.01		2.82		1.17	1.17	0.07	0.50	4 20	0.05	0.03
40	Aia & Dyar (2002)	MÖS	26.73	1.92	20.21		41 25	7.00	0.29	1 49	1 11		2.63	0 14	1.17	1 11	0.28	3.08	1.03	0.03	0.03
41	$\begin{array}{c} \text{Rillaut et al.} (2002) \\ \end{array}$	MÖS	37.02	1.72	35.61	0.89	41.23	11.07	0.07	1.49	1.11		3 23	0.14	0.77	3 11	0.03	0.02	1.65	0.05	1 19
42	Billaut <i>et al.</i> (2002)	MÖS	36 70		33.04	3 74		11.07	0.07	0.65			3 32		0.68	2.93	0.05	0.02	1.05		1.12
43	Billaut <i>et al.</i> (2002)	MÖS	36.81		30.58	7.25		8 13	0.11	1 14			3.42		0.58	2.95	0.33	0.15	1.20	0.01	1.20
44	Bertoldi <i>et al.</i> (2002)	MÖS	22.46	0.06	19 29	4 70	37 20	2.87	0.07	1.1.1		86 65	2.62	0.01	1 38	1.28	0.55	3.63	0.50	0.01	0.16
45	Bertoldi <i>et al.</i> (2001)	MÖS	30.38	0.00	18 70	0.79	6 2 9	29.81	0.05			86.12	2.94	0.01	1.06	1.08	0.06	0.51	4 30	0.00	0.04
46	Lougear $et al.$ (2000)	MÖS	00.00	0.10	101/0	0.7.2	0.2	2,101	0.02			00.12	2.68	0.01	1.32	1.18	0.18	1.95	2.68	0.00	0.02
47	Mitra & Bidvananda (2001)	MÖS	27.69	0.02	22.26		12.26	24.56	0.17	0.01	0.08	87.05	2.73	0.00	1.27	1.32	0.37	0.64	3.61	0.01	0.22
48	Mitra & Bidvananda (2001)	MÖS	26.74	0.06	21.77		11.72	24.36	0.17	0.01	0.17	85.00	2.70	0.01	1.29	1.31	0.37	0.62	3.67	0.02	0.19
49	Mitra & Bidyananda (2001)	MÖS	27.42	0.05	22.11		12.28	24.72	0.15	0.01	0.16	85.67	2.72	0.00	1.28	1.30	0.38	0.64	3.65	0.01	0.20
50	Smyth <i>et al.</i> (1997)	MÖS	31.68	0.05	17.13	1.34	1.62	34.76	0.08		0.05		2.99	0.00	1.01	0.89	0.10	0.13	4.88	0.01	0
51	Malmström et al. (1996)	MÖS											3.50		0.50	0.70	0.10	0.10	4.90		0.15
52	May et al. (1995)												3.00		1.00	0.90	0.10	0.10	4.90		0
53	Gregori & Mercader (1994)	CHE											3.12		0.88	1.69	0.64	1.03	2.17		0.73
54	Gregori & Mercader (1994)	CHE											3.01		0.99	0.49	0.29	0.50	4.20		0
55	Gregori & Mercader (1994)	CHE											2.76		1.24	2.01	0.12	0.70	3.30		0.45
56	Pal et al. (1993)	MÖS	27.69	0.02	22.26		12.26	24.56	0.17	0.01	0.08	87.05	2.73	0.00	1.27	1.32	0.38	0.63	3.61	0.01	0.21
57	Pal et al. (1993)	MÖS	26.74	0.06	21.77		11.72	24.36	0.17	0.01	0.17	85.00	2.70	0.00	1.29	1.30	0.37	0.62	3.67	0.01	0.19
58	Pal et al. (1993)	MÖS	27.42	0.05	22.11		12.28	24.72	0.15	0.01	0.16	86.90	2.71	0.00	1.29	1.30	0.38	0.64	3.65	0.01	0.19
59	Anceau (1992)	CHE	35.70		33.77	2.81	0.35	13.26	0.04		0.06	86.00	3.24		0.76	2.85	0.19	0.03	1.79		1.14
60	Kawano & Tomita (1991)		39.25	0.48	32.94	2.98		9.49		0.26	0.02		3.54	0.03	0.43	3.07	0.20		1.28		1.41
61	Chamberlain et al. (1989)		24.40		21.14		32.65	11.74	0.07				2.59		1.41	1.24	0.16	2.73	1.86	0.01	0
62	Joswig & Fuess (1990)												2.85		1.15	0.97	0.18	0.28	4.55		0.00
63	Zheng & Bailey (1989)												2.96		1.04	0.54	0.23	0.09	4.73		0
64	Ahn & Buseck (1988)												2.83		1.17	4.32	0.07				1.61
65	DeGrave et al. (1987)												2.88		1.12	1.32	0.23	2.07	2.16	0.05	0.22
66	DeGrave et al. (1987)												2.68		1.32	1.06	0.14	1.44	3.06	0.07	0
67	DeGrave et al. (1987)												2.64		1.36	0.85	0.32	1.82	3.10		0

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68	DeGrave et al. (1987)											2.69		1.31	1.55		0.86	3.47		0.12
69	Rule & Bailey (1987)											2.62		1.38	1.28	0.14	1.16	2.96		0.02
70	Ballet et al. (1985)											2.64		1.36	1.08	0.92	2.75	0.93		0.32
71	Ballet et al. (1985)											2.97	0.35	0.68	1.85	0.33	2.41	0.83		0.75
72	Ballet et al. (1985)											2.68		1.32	1.18	0.18	1.95	2.69		0.02
73	Ballet et al. (1985)											3.00		1.00	0.96	0.14	0.48	4.30		0.05
74	Lin and Bailey (1985)	32.90	0.06	35.40	2.69		14.73	0.01	0.01	0.07		3.00		1.00	2.81	0.18		2.00		1.00
75	MacKenzie et al. (1983)	20.95		35.21		8.28	22.88		0.58		88.90	2.01		1.99	2.00		0.66	3.28		0.00
76	MacKenzie et al. (1983)	21.42	0.26	34.06	0.63	0.95	28.66		0.51		86.49	2.02	0.02	1.98	1.82	0.04	0.08	4.04		0
77	Borggaard et al. (1982)											2.68		1.32	1.17	0.18	1.95	2.69		0.02
78	Kittrick (1982)											2.97		1.03	1.44	0.07	0.99	3.24		0.24
79	Kittrick (1982)											2.97		1.03	1.44	0.07	0.99	3.24		0.24
80	Kittrick (1982)											2.99		1.01	1.39	0.21	0.57	3.52		0.30
81	Kittrick (1982)											2.47		1.53	1.60		3.29	1.05		0.04
82	Kittrick (1982)											2.84		1.16	1.75	0.12	2.61	1.16		0.36
83	Kodama et al. (1982)	33.55		22.62	2.51	15.74	25.58	_				2.91		1.09	1.23	0.16	1.14	3.31		0.15
84	Kodama et al. (1982)	25.89		19.85	12.61	30.57	4.77	6.31				2.61		1.39	0.96	0.95	2.57	0.72	0.54	0.26
85	Kodama et al. (1982)	30.66		22.69	14.58	25.48	5.50	1.09				2.89		1.11	1.42	1.04	2.01	0.77	0.09	0.67
86	Kramm (1980)	32.81	0.18	35.71	2.66	1.07	13.54	0.29	0.06			2.99	0.01	1.00	2.85	0.18	0.08	1.84	0.02	1.01
87	Kramm (1980)											3.03		0.97	2.86	0.14	0.14	1.89	0.03	1.01
88	Kramm (1980)											3.04		0.96	2.87	0.13	0.14	1.88	0.04	1.02
89	Phillips et al. (1980)	32.20	0.03	15.23		1.19	35.56			3.40		3.02		0.98	0.72	0.04	0.04	4.95		0
90	Phillips et al. (1980)	31.58	0.04	15.37		1.42	34.93			3.57		2.99		1.01	0.71	0.06	0.06	4.92		0
91	Goodman & Bain (1979)											2.85	0.09	1.07	1.26	0.08	0.26	4.24		0.14
92	Goodman & Bain (1979)											3.12		0.89	0.60	0.18	0.32	4.95		0
93	Goodman & Bain (1979)											2.81		1.20	1.01	3.02	0.64	0.79	0.33	1.42
94	Goodman & Bain (1979)											3.12		0.88	1.47	0.13	3.18	0.73	0.12	0.35
95	Fransolet & Bourguignon (1978)	33.06		35.74	2.74	0.24	14.09	0.28				3.01		0.99	2.84	0.19	0.02	1.91	0.02	1.02
96	Fransolet & Bourguignon (1978)	32.82		36.16	2.45	0.05	14.35	0.24				2.99		1.01	2.86	0.17		1.94	0.02	1.01
97	Sudo & Shimoda (1978)	34.36		38.34	0.49		10.21		0.25			3.12		0.88	3.29	0.03		1.39		1.16
98	Sudo & Shimoda (1978)	35.67	0.19	37.41	0.20	0.09	11.97	0.02	1.25			3.18	0.01	0.81	3.12	0.01	0.01	1.59		1.12
99	Alysheva et al. (1977)	33.71		37.95	2.74	2.40	12.08		1.49			2.97		1.03	2.90	0.18	0.17	1.58		1.03
100	Alysheva et al. (1977)	36.07		32.42	1.99		15.94		0.44			3.24		0.76	2.67	0.13		2.14		0.99
101	Alysheva et al. (1977)	33.83	0.34	31.80	0.27	0.37	18.86		2.42			3.04	0.03	0.93	2.43	0.03	0.04	2.50		0.77
102	Ericsson et al. (1977)											2.95		1.05	0.46	0.08	0.07	2.39		0
103	Ericsson et al. (1977)											2.82		1.19	0.55	0.53	1.59	2.60	0.04	0

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TABLE 2. (contd.)

Nº	Deference	Method	SiO	TiO	A1 0	Fe O	E ₂ O	MaO	MnO	$CaO + Na_2O + K_O$	Others	Total	c;	ті	IVA1	VIA1	Fe ³⁺	Fe ²⁺	Ma	Mn	
IN	Kelelellee	wiethou	5102	1102	AI ₂ O ₃	re ₂ 0 ₃	гео	MgO	MIIO	K ₂ 0	Others	Total	51	11	AI	AI	ге	ге	Mg	IVIII	
104	Malysheva et al. (1977)												2.54		1.46	0.90	1.13	2.60	1.10		0.28
105	Spanu et al. (1977)												2.49		1.51	1.51	0.61	2.53	0.71	0.25	0.31
106	Spanu et al. (1977)												2.54	0.05	1.41	1.05	0.54	2.19	2.00	0.12	0.09
107	Spanu et al. (1977)												2.55		1.45	1.37	0.33	1.50	2.65	0.03	0.13
108	Tsukuki & Honda (1977)												3.19		0.81	3.13	0.01	0.01	1.60		1.13
109	Honda (1975)		34.36		38.84	0.49		10.21		0.85		85.75	3.12		0.88	3.29	0.03		1.38		1.16
110	Vrublevskaya et al. (1975)		38.26		44.28	1.31	0.48				2.00		3.42		0.58	4.08	0.09	0.04			1.79
111	Kimbara & Nagata (1974)		33.10	0.30	37.43	0.87		13.12		0.44		85.26	3.01	0.02	0.97	3.05	0.06		1.78		1.04
112	Kimbara & Sudo (1973)		29.45	0.20	20.14	4.08	18.69	14.93	0.36	1.59		89.34	2.97	0.02	1.02	1.36	0.31	1.58	2.24	0.03	0.27
113	Kimbara et al. (1973)		30.84	0.12	14.15	2.59	18.42	20.62	0.10	0.84		87.68	3.15	0.01	0.84	0.87	0.20	1.58	3.14	0.01	0.09
114	Post & Plummer (1972)		25.80		18.30	4.10	21.30	20.30				89.80	2.64		1.36	0.85	0.32	1.82	3.10		0
115	Post & Plummer (1972)		27.90		25.20		10.60	24.20				87.90	2.69		1.31	1.55		0.85	3.48		0.12
116	Post & Plummer (1972)		30.50		17.30	2.90	6.10	32.80				89.60	2.87		1.13	0.79	0.21	0.48	4.60		0
117	Post & Plummer (1972)		33.90		14.60	2.70	2.10	34.60				87.90	3.16		0.84	0.77	0.19	0.16	4.81		0.06
118	Aleksandrova et al. (1972)												3.14		0.86	4.10	0.04	0.01	0.08		1.64
119	deWaal (1970)		27.27		15.21	4.35	2.78	10.13	0.06	0.28	29.87	89.98	2.99		1.01	0.95	0.36	0.25	1.65	0.01	0.15
120	de Waal (1970)		27.27		15.21	4.35	2.78	10.13	0.06	0.38	29.87		2.98		1.02	0.95	0.36	0.25	1.65	0.01	0.14
121	Shimoda (1970)		32.72	0.03	20.28	1.00		24.47	0.06	2.22		80.78	3.24	0.00	0.76	1.61	0.07		3.61	0.01	0.44
122	Drits & Lazarenko (1967)		31.17		37.49	1.04	1.95	14.25		0.26			2.85		1.15	2.90	0.07	0.15	1.95		0.91
123	Gomes (1967)												3.45		0.55	3.96	0.02	0.02	0.04	0.01	1.72
124	Weaver et al. (1967)												2.66		1.34	1.74	0.02	3.68	0.42		0.21
125	Frenzel & Sehembra (1965)												3.43		0.57	2.70	0.09	0.04	1.97		1.11
126	Horikoshi (1965)		27.31	0.22	19.69	3.23	18.58	19.01	0.27	0.80		89.11	2.78	0.02	1.20	1.16	0.25	1.58	2.88	0.02	0.11
127	Hayashi & Oinuma (1964)		35.63		34.87	5.01	0.43	8.63	0.05	1.83		86.45	3.26		0.74	3.02	0.34	0.04	1.17		1.31
128	Tsukahara (1964)		39.01	0.47	32.15	0.90	0.10	10.14		2.16			3.56	0.03	0.41	3.05	0.06	0.01	1.37		1.25
129	Tsukahara (1964)		30.66		30.55	0.47		10.13		1.79			3.24		0.76	3.04	0.04		1.59		1.00
130	Banno (1964)		24.45	0.21	21.99	20.80	28.04	11.90	0.54	0.25		89.46	2.59	0.02	1.39	1.36	0.17	2.19	1.88	0.05	0.06
131	Caillère et al. (1962)												2.70		1.30	4.30	0.03	0.02			1.52
132	Brydon et al. (1961)	CHE											2.80	0.09	1.11	3.31	0.38		0.76		1.18
133	Bilgrami & Howie (1960)		28.32	0.09	19.03	1.19	14.85	23.72	0.09	0.63		87.92	2.83	0.01	1.17	1.06	0.09	1.24	3.53	0.01	0
134	Nakamura (1960)		21.34	0.21	20.73	5.20	40.17	0.99	0.89	0.04		89.57	2.46	0.02	1.52	1.30	0.45	3.87	0.17	0.09	0.12
135	Lapham (1958)		36.43		12.24	0.94	6.87	30.94	0.11	0.33	0.63	88.49	3.43		0.57	0.79	0.07	0.54	4.35	0.01	0.15
136	Shimane & Sudo (1958)		33.14	0.08	10.04	4.00	5.66	32.64		0.80		86.36	3.25	0.01	0.75	0.41	0.29	0.46	4.77		0
137	Shirozu (1958)		29.73		17.95	0.68	10.05	21.84	8.24			88.49	2.98		1.02	1.09	0.05	0.84	3.26	0.70	0.06
138	Shirozu (1958)		22.81		20.35	4.84	34.87	6.11	0.56	0.10			2.54		1.46	1.20	0.40	3.24	1.01	0.05	0.07

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139	Shirozu (1958)	23.38		19.32	2.51	38.38	5.39	0.58	
140	Shirozu (1958)	26.04		19.96	1.85	21.34	18.56	0.47	
141	Shirozu (1958)	29.71		19.75	0.29	6.59	30.79	0.30	
142	Shirozu (1958)	32.27		14.99	1.56	6.89	31.75	0.09	
143	Battey (1956)	27.11	0.35	17.42	2.91	30.98	9.75		0.21
144	Frondel (1955)	29.94	0.16	13.77	10.46	20.96	10.71	0.18	1.34
145	Stone & Weiss (1955)	28.73	0.41	19.16	1.97	10.99	26.37	0.15	0.10
146	Brindley & Youell (1953)	22.47		21.82	0.22	37.24	2.57		
147	Brindley & Youell (1953)	23.48		22.45	43.37	0.45	2.60		
148	Ginzburg (1953)	33.40		47.47		0.71	0.20		0.54
149	Mathias (1952)	29.84	0.15	20.20	0.80	15.00	21.83	0.13	0.80
150	Brindley (1951)	22.03	3.63	22.91	0.46	36.68	1.91	0.04	0.18
151	Brindley (1951)								
152	Brindley (1951)								
153	Brindley (1951)	21.40	0.89	25.40	0.25	37.60	2.04	0.05	0.13
154	Brindley (1951)								
155	Bannister & Whittard (1945)	26.40		18.23	5.70	25.87	11.35	0.04	0.76
156	Hutton & Seelye (1945)	27.64	0.22	22.48	0.06	12.06	24.32	0.02	0.23
157	Hutton & Seelye (1945)	31.87	0.17	14.51	1.86	3.57	32.76		0.06
158	Engelhardt (1942)	20.82		17.64	8.70	37.96	4.15		
159	Engelhardt (1942)	26.65		16.14	6.69	34.43	4.47		
160	Hödl (1942)	22.47		23.57	4.01	29.27	9.81		
161	Hutton (1940)	26.45		20.88	2.82	21.06	16.84	0.44	0.47
162	Lazarenko (1940)								
163	Lazarenko (1940)								
164	Lazarenko (1940)								
165	Hallimond et al. (1939)								
166	Hallimond et al. (1939)	24.35	0.04	20.21	2.13	36.27	5.57	0.48	0.10
167	Hallimond et al. (1939)	24.90		15.60	7.20	35.00	4.60	0.40	
168	Nockolds & Richey (1939)	25.07	0.12	19.78	3.50	35.80	1.11	0.50	2.15
169	Holzner (1937)	25.09	0.05	18.94	9.60	25.66	9.80	0.02	0.49
170	Holzner (1937)	22.18	0.04	20.04	7.35	35.23	3.79	0.02	0.49
171	Hutton (1938)	26.69	0.30	19.57	3.49	21.80	16.23	0.30	0.17
172	Mélon (1938)	27.56		24.47	3.85	10.51	20.86	1.80	
173	Tilley (1938)	25.62	0.88	21.19	3.88	21.55	15.28	0.35	0.16
174	Simpson (1937)	23.32	0.03	17.45	4.09	38.90	4.54	0.01	0.24
175	Simpson (1936)	23.43		18.21	4.49	36.30	5.41		
176	Ross (1935)	24.60		19.10	3.12	29.04	13.10		

	89.56	2.63		1.37	1.20	0.21	3.61	0.90	0.06	0.02
	88.22	2.69		1.31	1.13	0.14	1.85	2.86	0.04	0
	87.43	2.84		1.16	1.07	0.02	0.53	4.39	0.02	0
	87.55	3.10		0.90	0.79	0.11	0.55	4.54	0.01	0.00
	88.73	2.93	0.03	1.07	1.15	0.24	2.80	1.57		0.16
	87.52	3.19	0.01	0.81	0.92	0.84	1.87	1.70	0.02	0.45
	87.88	2.81	0.03	1.19	1.03	0.15	0.90	3.85	0.01	0
	84.32	2.65		1.35	1.68	0.02	3.67	0.45		0.18
	92.35	2.35		1.65	1.00	3.27	0.04	0.39		1.31
3.12	85.44	3.07		0.93	4.22		0.05	0.03		1.64
0.03	88.78	2.93	0.01	1.07	1.27	0.06	1.23	3.19	0.01	0.12
0.26	88.77	2.47	0.31	1.53	1.50	0.04	3.44	0.32	0.00	0
		3.10		0.90	1.30	0.60	3.35	0.75		0.50
		2.48	0.30	1.22	1.82	0.04	3.46	0.32		0.32
	87.98	2.42	0.08	1.58	1.80	0.02	3.55	0.34	0.00	0.12
		2.48	0.30	1.22	1.82	0.04	3.46	0.32		0.32
	88.35	2.83		1.17	1.13	0.46	2.32	1.81	0.00	0.18
0.06	87.09	2.73	0.02	1.27	1.34	0.00	0.99	3.57	0.00	0.01
1.38	86.18	3.07	0.01	0.93	0.72	0.14	0.29	4.71		0
	89.27	2.42		1.58	0.83	0.76	3.69	0.72		0.00
	88.38	2.99		1.01	1.12	0.56	3.23	0.75		0.34
	89.13	2.42		1.58	1.40	0.32	2.63	1.57		0.07
	88.96	2.71		1.29	1.24	0.22	1.81	2.57	0.04	0.06
		3.09		0.91	3.92	0.05		0.21		1.44
		3.06		0.94	3.88	0.08		0.23		1.40
		3.10		0.90	4.16	0.07		0.11		1.67
		2.80		1.20	0.86	0.60	3.56	0.78	0.04	0.13
	89.15	2.70	0.00	1.30	1.35	0.18	3.37	0.92	0.05	0.12
	87.70	2.86		1.14	0.97	0.62	3.36	0.79	0.04	0.22
	88.03	2.85	0.01	1.15	1.50	0.30	3.40	0.19	0.05	0.24
	89.65	2.68	0.00	1.32	1.06	0.77	2.29	1.56	0.00	0.24
	90.60	2.51	0.00	1.49	1.18	0.63	3.33	0.64	0.00	0.15
0.13	88.68	2.76	0.02	1.24	1.15	0.27	1.89	2.50	0.03	0.09
	89.05	2.68		1.32	1.48	0.28	0.85	3.02	0.15	0.22
	88.91	2.64	0.07	1.36	1.22	0.30	1.86	2.35	0.03	0.08
	88.58	2.68	0.00	1.32	1.05	0.35	3.74	0.78	0.00	0.04
	87.84	2.68		1.32	1.13	0.39	3.47	0.92		0.10
	88.96	2.64		1.36	1.06	0.25	2.61	2.10		0

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N°	Reference	Method	SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	FeO	MgO	MnO	$\begin{array}{c} \mathrm{CaO} + \mathrm{Na_2O} + \\ \mathrm{K_2O} \end{array}$	Others	Total	Si	Ti	^{IV} Al	^{VI} Al	Fe ³⁺	Fe ²⁺	Mg	Mn	
177	Jung (1931)		26.65		16.14	6.69	34.43	4.47				88.38	2.99		1.01	1.12	0.56	3.23	0.75		0.34
178	Dschang (1931)		30.30		18.25	2.64	3.98	32.66		0.03		88.24	2.87		1.13	0.90	0.19	0.32	4.61		0
179	Jung & Kohler (1930)		20.82		17.64	8.70	37.96	4.15				89.27	2.42		1.58	0.83	0.76	3.69	0.72		0.00
180	Orcel, 1927, p246		22.27	0.08	21.40	0.67	43.01	2.35	0.05	0.15		90.33	2.54	0.01	1.46	1.41	0.06	4.10	0.40	0.00	0.00
181	Orcel (1927)		26.50	0.03	20.85	1.90	18.73	19.85	0.52			88.38	2.69	0.00	1.31	1.19	0.15	1.59	3.01	0.04	0.01
182	Orcel (1927)		25.35		22.06	18.97	8.36	15.06	0.08			89.88	2.52		1.48	1.11	1.42	0.70	2.24	0.01	0.53
183	Orcel (1927)		27.12		27.68	0.20	1.24	30.96	0.54			87.74	2.51		1.49	1.54	0.01	0.10	4.28	0.04	0.03
184	Chernykh (1926)		25.04		20.10	2.05	35.40	4.28				86.87	2.83		1.17	1.51	0.17	3.34	0.72		0.25
185	Shannon & Wherry (1922)		23.20		24.42	3.48	13.40	32.76		1.04		88.30	2.10		1.90	0.71	0.24	1.01	4.42		0
186	Shannon & Wherry (1922)		27.78		24.30	1.43	0.35	32.71				86.57	2.61		1.39	1.29	0.10	0.03	4.58		0.00
187	Poitevin & Graham (1918)		24.40		22.77	0.45		32.70	0.09	0.40		80.81	2.47		1.53	1.18	0.03		4.93	0.01	0
188	Dana (1915)		33.24		11.07	2.26	25.11	16.51	0.41	1.36		89.96	3.42		0.58	0.77	0.18	2.16	2.53	0.04	0.16
189	Tschermak (1891)		23.62		22.26		38.97	1.09	0.98	1.67		88.59	2.68		1.32	1.67		3.70	0.18	0.09	0.03
190	Tschermak (1891)		33.83		12.95	2.25	3.02	34.94				86.99	3.21		0.79	0.65	0.16	0.24	4.94		0.01
191	Hawes (1875)		33.46		10.96	2.56	24.72	16.52	0.40	1.21		89.83	3.44		0.56	0.77	0.20	2.13	2.53	0.03	0.18

Footnote: Speciation of Fe was obtained using chemical titration (CHE) or, for the past 20 years, using XANES (XAN), Mössbauer (MÖS) or XPS spectroscopies. Where no method is indicated, chemical titration is assumed.



FIG. 1. Backscattered-electron image of an oscillatoryzoned chlorite showing Al-free di-trioctahedral substitution, from the Pic-de-Port-Vieux, Pyrenees, Spain. The numbers refer to the structural formulae reported in Table 2.

rather than by changes in the buffering assemblage (Lanari *et al.* 2014a; Trincal *et al.*, 2015). Most of the spot analyses reported in the literature correspond to different geological environments (Table 2); it is thus very difficult to use them to evaluate the link between the Fe³⁺ behaviour and the evolution of a parameter such as pressure, temperature, f_{Ω} , *etc.*

The choice of analytical method to estimate the XFe^{3+} in chlorite is essential. Chlorite Fe³⁺ measurements by Mössbauer spectroscopy or titration are relatively common but provide few constraints due to poor spatial resolution. Indeed, recent investigations using high-resolution mapping (*e.g.* de Andrade *et al.*, 2006; Muñoz *et al.*, 2006) have demonstrated that chlorite is frequently zoned, recording strong compositional variability even at the crystal scale (*e.g.* Lanari *et al.*, 2014b). Compositional zoning reflects time series of equilibrium conditions experienced by the rock. Thus *in situ* micrometric analyses are required to ensure a precise determination of the Fe³⁺/Fe_{Tot} of the successive growth zones, which is impossible using Mössbauer spectroscopy. μ -XANES *X*Fe³⁺ spot analyses or mapping allow us to elucidate the heterogeneity of the Fe chemistry of chlorite which is critical to see the relations between formation temperature and crystal geometry (Vidal *et al.*, 2006; Trincal *et al.*, 2015).

µXANES DATA FROM ZONED CHLORITES

Accurate in situ µXANES and electron microprobe analyses performed on zoned chlorite from the Sambagawa metamorphic belt in Japan (Vidal et al., 2006) and on oscillatory-zoned chlorite from hydrothermal veins in the Pic-de-Port-Vieux, Pyrenees, Spain (fig. 1 from Trincal et al., 2015) show variations in the composition of the Fe^{3+} , Mg and Fe^{2+} cation proportions, while that of Si and Al remain constant (Table 2, Fig. 2). This trend is observed independently in both samples, which have different bulk-rock compositions and experienced different metamorphic conditions. Indeed, chlorite from Japan comes from a blueschist-facies metapelite made of K-white mica, garnet, albite, quartz and graphite (Vidal et al., 2006), while the chlorite from the Pyrenees crystallizes in a vein of a greenschist-facies metapelite without garnet or graphite (Trincal et al., 2015). The absence of graphite is a good indicator of more oxidizing conditions. In both cases, XFe³⁺ increases as temperature decreases, as suggested by numerous authors (e.g. Lanari et al., 2014a; Vidal et al., 2016). Furthermore, in both cases, there is no significant change in the Al content of the zoned chlorite. This result excludes the hypothesis of DT, TK or Al-Fe³⁺ substitutions being at the origin of the XFe^{3+} variations.

AL-FREE DI-TRIOCTAHEDRAL SUBSTITUTION

Based on μ XANES analyses performed on zoned chlorite by Vidal *et al.* (2006) and Trincal *et al.* (2015), and in agreement with the literature data (Table 2), we demonstrate that a di-trioctahedral substitution implying Fe³⁺ can be used to explain the variations of *X*Fe³⁺ in natural chlorite:

$$2Fe^{3+} + \Box = 3(Mg^{2+}, Fe^{2+})$$
(1)

As this substitution does not explicitly involve any Al cations, it is named Al-free di-trioctahedral substitution (AFDT); it must not be confused with the Al-free chlorite end-member.

This substitution is similar to another couple suggested by Billault *et al.* (2002) and based on electron microprobe



FIG. 2. Compilation of the Fe³⁺-bearing chlorite compositions. The Si-Al plot is used to highlight the di-trioctahedral (DT) and Tschermak (TK) substitutions. End-members and structural formulae are reported in Tables 1 and 2.



FIG. 3. Compilation of Fe^{3+} -bearing chlorite compositions. The $\text{Si}^{4+}-R^{2+}$ plot by Wiewióra & Weiss (1990) is used to highlight the di-trioctahedral (DT) and the Al-Fe³⁺ substitutions.



FIG. 4. Compilation of the Fe^{3+} -bearing chlorite compositions. The $\text{Fe}^{3+}-R^{2+}$ plot highlights the di-trioctahedral (DT), Al-Fe³⁺ and AFDT substitutions.

and Mössbauer spectroscopic analyses on sudoite:

$$Fe^{3+} = Al_{(VI)}$$
 and $Fe^{2+} + Fe^{3+} + \Box = 3(Mg^{2+})$

(2)

However, chlorite analyses from their study display both Al and XMg variations, which is not the case for the samples in the present study.

Chlorite compositions from Vidal *et al.* (2006) and Trincal *et al.* (2015) are plotted in an Fe³⁺ *vs. R*²⁺ diagram (Fig. 4). In this diagram, data from zoned chlorite are distributed along a line parallel to the AFDT substitution vector. The AFDT substitution allows us to define a theoretical Fe³⁺-bearing chlorite end-member; namely ferri-sudoite ((Si₃A1)[(Fe²⁺, Mg)₂Fe³⁺_A1)O₁₀(OH)₈) with two atoms of Fe³⁺ (Table 1). Ferri-sudoite can be separated into ferri-sudoite-Fe ((Si₃A1)[Fe²⁺Fe³⁺_2A1)O₁₀(OH)₈) and ferri-sudoite-Mg ((Si₃A1)[Mg₂Fe³⁺_2A1)O₁₀(OH)₈). As Fe³⁺ replaces Mg and Fe²⁺, it seems convenient to allocate up two atoms of Fe³⁺ to the M2-M3 sites (total multiplicity of 4) leaving Al in the M4 (Table 1).

According to several authors, the Fe³⁺ cation seems limited to 1 a.p.f.u and is generally allocated to the M4 site following the Al-Fe³⁺ substitution. This last assertion is supported by Z-contrast images of highangle annular dark-field imaging-scanning transmission electron microscopy (HAADF-STEM) and other studies (*e.g.* Smyth *et al.*, 1997; Inoue & Kogure, 2016; Vidal *et al.*, 2016). Among data collected in the literature, only six analyses show Fe³ ⁺ contents of >1 a.p.f.u (Orcel, 1927; Brindley & Youell, 1953; Malysheva *et al.*, 1977; Goodman & Bain, 1979; Kodama *et al.*, 1982; Singer *et al.*, 2009 in Table 2). However, these analyses would actually match other phyllosilicates, suggesting contamination issues. Although AFDT substitution suggests that it is possible, no sudoite with >1 Fe³⁺ atom has been yet discovered; it offers interesting research opportunities.

CONCLUSION

The compilation of Fe³⁺-bearing chlorite analyses in this study, especially those acquired by μ XANES spectroscopy in zoned crystals, suggests the existence of a new chlorite substitution, the AFDT substitution that is characterized by $2Fe^{3+} + \Box = 3(Mg^{2+}, Fe^{2+})$ and implies a new chlorite end-member: ferri-sudoite $(Si_3A1)[(Fe^{2+},Mg)_2Fe_2^{3+}\Box A1)O_{10}(OH)_8$ with 2 atoms of Fe³⁺. These results are not in agreement with previous studies which showed Al-Fe³⁺ substitution implying only 1 Fe³⁺. The absence of constraints on the Fe³⁺ crystallographic position together with the Fe³⁺ ionic properties in chlorite require further studies to confirm

the existence of the AFDT substitution and of the ferrisudoite end-member, *e.g.* with accurate *in situ* Fe^{3+} measurements in synthetic or diagenetic chlorite formed in different oxidizing environments. This study opens the door for further research developments that would have strong implications, *e.g.* in chlorite geothermometry.

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